

# Thermal infrared data of active lava surfaces using a newly-developed camera system

# Introduction

A field-portable miniature multispectral thermal infrared camera (MMT-Cam) was developed as part of the January 2017 HyspIRI Preparatory project. The goal of the project is to assess the research applications that are possible with thermal infrared (TIR) data, at spatial and spectral resolutions proposed by this mission. The MMT-Cam was built to acquire accurate emissivity data in situ with changing temperature on active lava surfaces. Furthermore, constraining the relationship between the emissivity spectral change and radiance derived from TIR data will provide more accurate temperatures<sup>[1]</sup>. Collection of accurate temperature and emissivity data during lava flow emplacement will greatly improve models designed to predict flow dynamics and down-flow hazard potential<sup>[1]</sup>. Finally, through spatial degradation analysis, constraints can be improved for the identification of changes in temperature and emissivity during cooling at lower spatial resolutions.

## Location

The first field campaign was conducted at Kilauea volcano, Hawai'i in January 2017 (Fig. 1a). Kīlauea volcano is a basaltic shield volcano located on the eastern slope of Mauna Loa volcano<sup>[2]</sup>. During the campaign two volcanic processes were targeted:

- I. Lava flows (primary) the surface lava flow activity on the pali and coastal plains on the eastern slopes of Kilauea volcano produced by the episode 61g flows (July 2016 - present) from Pu'u 'Ō'ō vent (Fig. 1b)<sup>[3]</sup>.
- 2. Lava Lake (secondary) the 250 m long and 190 m wide active lava lake within the Halema`uma`u crater<sup>[4]</sup> (Fig. 1c).

Additionally, further testing was undertaken at the lava pour experimental facility run by the Syracuse University Lava Project team, where melts of basaltic compositions are poured onto a sand substrate (Fig. 1d).



Figure 1: a) The island of Hawai'i, white arrows showing field areas (source: ESRI). b) MMT-Cam deployed at the lava flow entering the ocean near Kamokuna and c) Halema`uma`u lava lake. d) Lava pour at Syracuse University.





## Instrumentation



Specification	MMT-Cam	Figure 2: a		
Core	FLIR A65 (2 <sup>nd</sup> generation)	window, <b>b)</b> i showing the the MMT-Ca		
Spatial resolution	640 x 512 pixels			
Field of view (FOV)	45° x 37° with 13 mm lens			
Image frequency	30 Hz			
Gain settings	-25°C to 135°C / -40°C to 550°C	Table 1. Tec		
Detector	Uncooled VOX microbolometer	new MMT-Ca		
Spectral resolution	7.5 – 13.0 μm			
Filter Wheel	7 port – 6 IR filter + 1 open port			

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Band Number	Band Center [µm]	Peak Transmission [%]	Spect (FW
			Min
Broadband	11.00	91.5	7.500
1	8.04	88.6	7.814
2	8.55	89.1	8.313
3	8.99	84.1	8.766
4	9.55	90.0	9.320
5	10.04	86.7	9.805
6	11.35	73.5	11.086

## Calibration

The MMT-Cam system calibration incorporates only data where the camera and filters are perfectly aligned. Second, blackbody experiments were conducted at the Aerospace Corporation using environmentally controlled blackbodies at temperatures from 10 to 200 °C. The per-pixel radiance measured at the focal plane array (FPA) were compared to the actual filter and broadband blackbody radiance values (Fig. 4a). Second order polynomial models were produced to correct for any attenuations and errors associated with the system design.

Multispectral TIR data acquired during a lava pour experiment at Syracuse University show a change in emissivity spectra of basaltic lava surfaces as they cool. During emplacement and subsequent cooling from 1270 to 730 °K, the emissivity spectra (Fig. 5b) shallow and spectral contrast increases.



Figure 5: a) Calibrated broadband TIR image acquired during the lava pour experiment, the labels indicate the location where the emissivity spectra were extracted from. b) Emissivity spectra, calculated using the Emissivity Normalization method <sup>[5]</sup>, from the proximal molten channel of the flow (1270 °K) and the distal cooling lobe of the flow (730 °K).

The preliminary lava lake data (Fig. 6a) show that the primary emissivity absorption feature (around 8.5 to 9.0 µm) transitions to emissivity between 9 to 10 µm and shallows as the lava surface cools from 760 to 520 °K, forming a progressively thicker crust. The spectrum is likely a mixture of both the lava surface and SO<sub>2</sub>. The feature transitioning to longer wavelengths is partially due to the structural and physical state change as the glassy crust forms over the molten material (Fig. 6b). This change in the Si-O-Al bond structure is postulated to be changing the emissivity and with time, the spectrum should approach that of a solid cold basalt. This is the first time that accurate, unsaturated emissivity data with changing temperature has been measured in situ on active lava surfaces.



Figure 6: a) Before and after calibration TIR images of the Halema`uma`u crater lava lake acquired with the 8.04 µm filter. b) Emissivity difference spectra of molten and glass lava surface, acquired from the lava pour and lava lake <sup>[5]</sup>.

Future work will include applying the current methodology to evaluate the spatiotemporal variability in temperature and emissivity during natural lava flow emplacement and cooling. A correction for the influence of SO<sub>2</sub> on the emissivity spectra will be developed. The corrected spatiotemporal variability in emissivity of propagating lava flows will be incorporated into prediction models to improve outputs. Finally, these results and methodologies will be compared to current and proposed satellite datasets, including ASTER and HyspIRI, to determine future capabilities.

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[1] Ramsey, M.S. & Harris, A.J.L., 2013. Volcanology 2020: How will thermal remote sensing of volcanic surface activity evolve over the next decade? Journal of Volcanology and Geothermal Research, 249, pp.217–233. [2] Orr T, Poland MP, Patrick MR, Thelen WA, Sutton AJ, Elias T Thornber CR, Parcheta C, Wooten KM. 2015. Kilauea's 5-9 March 2011 Kamoamoa fissure eruption and its relation to 30+ years of activity from Pu'u 'Ō'ō. In: Carey R, Poland M, Cayol V, Weis D, (eds) Hawaiian Volcanism: From Source to Surface: Hoboken, New Jersey, Wiley, American Geophysical Union Geophysical Monograph 208, p. 393-420. [3] Patrick MR, Orr T, Fisher G, Trusdell F, Kauahikaua J. 2016. Thermal mapping of a pahoehoe lava flow, Kilauea Volcano. Journal of Volcanology and Geothermal Research, 332:71-87. [4] Patrick M., Orr T., Sutton A.J., Elias T. Swanson D. 2013. The first five years of Kilauea's summit eruption in Halema`uma`u Crater, 2008-2013. [5] Gillespie, A. R., Rokugawa, S., Hook, S. J., Matsunaga, T., and Kahle, A. B. 1999. Temperature/emissivity separation algorithm theoretical basis document, Version 2.4. Jet Propulsion Laboratory, Pasadena, 22 March 1999.



# **Results and Conclusions**

# **Future Work**

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# References